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高寒干旱区地表水与地下水水化学特征及转换关系: 以大通河流域为例

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(1.中国地质科学院水文地质环境地质研究所,石家庄 050061; 2.中国地质大学(北京)水资源与环境学院,北京 100083; 3.中国地质环境监测院,北京 100081; 4.中国地质调查局河北省地下水污染机理与修复重点实验室,石家庄 050061) 摘要:全球气候变暖和区域极端气候会导致高寒干旱山区平原降雨失衡、冰川融化和冻土退化,进而改变区域水文循环.其中,地表水与地下水转化关系是高寒干旱区水文循环气候响应的重要科学问题之一.以祁连山南麓大通河流域为研究区,基于119 组基本水化学参数和 38 组氘氧同位素数据,利用多元统计分析和同位素技术,研究了流域地表水与地下水的水化学特征及其相互转化过程.结果表明,流域地表水以 HCO₃-Mg·Ca 型水为主,受控于岩石风化作用;地下水以 HCO₃-Mg·Ca 型和 Cl·SO₄-Na型水为主,受岩石风化作用和蒸发浓缩共同控制.上游地下水存在少量钙、镁长石的溶解,中游地下水化学组分主要为碳酸盐岩的风化溶解,下游地下水中各离子组分蒸发富集使地下水溶解性总固体升高.风化溶滤、人为活动、原生沉积环境、阳离子交替吸附及其他因素对研究区地表水和地下水化学组分的贡献率依次为 39.1%、15.0%、12.6%、13.8% 和19.5%. δD 和 δ¹δ0 同位素测试结果表明,沿地下水流向大通河河水氘氧同位素含量呈富集到贫化的变化趋势,大通河上、中游地区河水主要受大气降水补给,而下游河段受地质构造和水文地质条件等影响,主要为潜水和泉水溢出补给河水,为地下水排泄区.

关键词:地表水;地下水;水化学特征;氢氧稳定同位素;转化关系;大通河流域中图分类号: X143 文献标识码: A 文章编号: 0250-3301(2023)02-0752-09 **DOI**: 10.13227/j. hjkx. 202204232

Hydrochemical and Isotopic Evidence for Groundwater Conversion of Surface Water in Alpine Arid Areas: A Case Study of the Datong River Basin

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Abstract: Global warming and regional extreme climates will lead to unbalanced rainfall, melting glaciers, and permafrost degradation in alpine and arid mountain plains, thereby changing the regional hydrological cycle. The relationship between surface water and groundwater conversion is one of the important scientific issues of hydrological cycle climate response in alpine arid areas. Taking the Datong River Basin at the southern foot of the Qilian Mountains as the study area, based on 119 sets of basic hydrochemical parameters and deuterium-oxygen isotope data, using multivariate statistical analysis and isotopic techniques, the hydrochemical characteristics of surface water and groundwater in the basin and their mutual transformation process were studied. The results showed that the surface water was HCO₃-Mg·Ca type, which was mainly controlled by rock weathering, whereas the groundwater was HCO₃-Mg·Ca type and Cl·SO₄-Na type, which was controlled by rock weathering and evaporation concentration. There was a small amount of calcium and magnesium feldspar dissolved in the upstream groundwater, and the chemical components of the midstream groundwater were mainly the weathering and dissolution of carbonate rocks. The contribution rates of weathering filtration, anthropogenic activities, native sedimentary environment, alternating adsorption of cations, and other factors to the chemical components of surface water and groundwater in the study area were 39.1%, 15.0%, 12.6%, 13.8%, and 19.5%, respectively. The deuterium and oxygen isotope contents of Datong River water showed a trend of enrichment to depletion along the groundwater flow direction. The δD and δ¹⁸O isotope test results showed that the deuterium and oxygen isotope content in the Datong River along the groundwater flow showed a trend of enrichment to depletion. The upper and middle reaches of the Datong River were mainly recharged by atmospheric precipitation, whereas the lower reaches were affected by geological structure and influenced by h

Key words: surface water; groundwater; hydrochemical characteristics; oxygen and hydrogen isotopes; transformation relationship; Datong River Basin

近年来,全球气候变暖和区域极端气候导致中国西部高寒干旱地区山区降雨量增多而平原区减少^[1]、冰川融化^[2,3]和冻土退化^[4,5]等,进而改变水文循环过程和水资源量^[6]. 地表水与地下水的转化是流域水循环过程的重要组成部分,影响水资源量和水化学组分的时空分布与演化. 因此,研究极端气候下地表水与地下水的转化对于改善高原水资源可持续性和生态系统健康具有重要意义^[7~9].

基于传统基本水化学参数记录水体来源、运移和转化,是目前国内外用于判定地表水与地下水转化关系的重要技术手段之一,然而,水文地质条件^[10]、河流特性^[11]、季节性降水^[12]和人类活动^[13]

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与水质演化,E-mail: 793833881@qq.com * 通信作者,E-mail:2767398591@qq.com 等因素会影响基本水化学参数的指示精度,而且传统方法在定量溯源方面存在不足.为了解决上述问题,能够反映地表水与地下水相互转化量化关系的同位素技术被广泛应用^[14].苏小四等^[15]、雷义珍等^[16]和聂振龙等^[17]分别对马莲河、青海湖沙柳河和黑河流域河水和地下水的水化学和同位素特征进行了分析,以确定二者的转化关系.马莲河中游河水和地下水之间存在着微弱的水力联系;海湖沙柳河的 δD 和δ¹⁸O值在消融期最大,而在冰冻期最小;黑河流域位于祁连山区,主要受地下水补给,而在祁连山南麓山前戈壁,则是由出山河水人渗转换成地下水.刘双等^[18]对河流与地下水的δ¹⁸O值和水化学特征进行了定量分析,确定了三江平原挠力河干流和地下水转换量.

本文以我国典型高寒干旱区祁连山南麓大通河流域为研究区,基于119组基本水化学参数和38组水氘氧同位素数据,利用Piper图、Gibbs图和同位素技术,并结合多元统计分析方法对流域地表水与地下水的补排关系和相互转过程进行了研究,通过揭示高寒区河流水循环过程,以期为水资源的合理利用提供科学支撑.

1 研究区概况

大通河是黄河水系——湟水的一级支流,地处

祁连山南麓. 位于北纬 36°30′~38°25′和东经98°30′~103°15′之间,海拔为1 592. 58~4 420. 79 m,全长为 560. 7 km,流域面积为15 130 km²[19]. 研究区属大陆高原性气候,多年平均气温为 3. 3~8. 1℃,多年平均降水量为 494. 1 mm^[20],多年平均径流量为 28. 95×10⁸ m³[21],流域内水资源丰富,已建纳子峡和石头峡两座大型水库和"引大入秦"、"引硫济金"和"引大济湟"这 3 项大型调水工程^[22],水库和调水工程对地表径流、地下水动力场和化学场均有较大影响,导致生态环境及地质环境恶化^[23].

大通河流域以尕大滩、连城为界,划分为上、中和下游这3个区段^[23],见图1,河水和地下水流向均为自西北向东南,上游河岸两侧山地分布着碎屑岩类裂隙水,大气降水和冰雪融水形成基岩裂隙水或补给河水,此外,河谷和山前平原有第四系松散孔隙水分布,粗颗粒的冲洪积物为水循环提供良好的通道,使地表水和地下水转化频繁;流域中游海拔3800m以下非冻土区,以基岩裂隙水为主,因受大气降水和少量冻结层上水补给,含水岩组为三叠系及古生界、下元古界变质岩,侵入岩,富水性极不均匀;下游低山区分布着碎屑岩裂隙潜水和变质岩裂隙水,部分丘陵台地有风化裂隙水,泉水流量小于0.5 L·s^{-1[19]}.该流域是一个半干旱地区,生态系统脆弱,区内地表水和地下水转化频繁^[23].

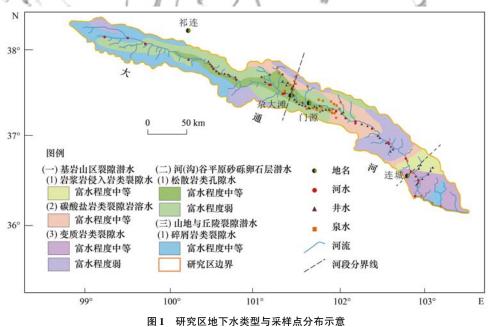


Fig. 1 Distribution map of groundwater types and sampling sites in the study area

2 材料与方法

2.1 样品采集与测试

2021年6~8月于大通河流域共采集样品119组.其中,地表水27组,浅层地下水59组,泉水33

组(图1). 大通河流域共采集河水、地下水稳定同位素水样 38 组(其中河水样 10 组,地下水样 28 组),分别用 550 mL和 2 L 的塑料瓶收集了水化学法和氘氧同位素分别法测定的样品,并在取样前用 0.45 μm 的滤膜进行过滤. 用 3 mol·L⁻¹硝酸酸化阳

离子至 pH < 2, 采集的样品用 Parafilm 封口膜密封后放置于 4~5℃环境中. 采用便携式野外测量仪 (HACH DR2800) 现场测量 pH、DO、水温和电导率等指标.

室内实验执行标准参照文献 [24],用电感耦合等离子体发射光谱仪(ICP-MS, Varian820-ms,美国)对水样的阳离子(K^+ 、 Na^+ 、 Ca^{2+} 和 Mg^{2+})进行测定.使用离子色谱仪(ISC1500)对阴离子(SO_4^{2-} 、 NO_3^{-})测定. NO_2^{-} 和 NH_4^+ 分别用紫外分光光度法和纳氏试剂光度法测定. Cl^- 、 HCO_3^- 和总硬度分别采用 $AgNO_3$ 滴定法、酸标准溶液滴定法和乙二胺四乙酸二钠滴定法测定. TDS 采用烘干测量法. 水样中 δ^{18} O和 δD 由水同位素分析仪(L2130i,皮卡罗,美国)测定,精度分别为±0.1‰和±0.8‰,测定结果均经过 Vienna Standard Mean Ocean Water (V-SMOW)转化.

2.2 数据分析

利用 Piper 图法可以有效判定水体中水化学类型及各离子含量占比^[25],故用 Piper 图对各河段水化学类型进行对比分析.主成分分析(PCA)是一种有效简化数据而不丢失重要信息的方法,可以帮助识别水质变量之间的关联,找到相同的来源并对水质参数进行分组^[26,27].若成分矩阵中 PC 为正相关,通常表示相同来源或相似的地球化学行为^[28,29].因此,本文采用主成分分析(PCA)对水化学参数分类以确定地表水、地下水的来源及贡献率.利用稳定氘氧同位素方法,可以分析河流径流的来源、河流径流的分布、地表水与地下水的转化.采用孙从建等^[29]提出的祁连山大气降水方程作为研究区大气降水线.根据同位素质量守恒和地表水补给来源的差异,利用二端元混合模型计算了各水源对地表水量的补给比例^[31].

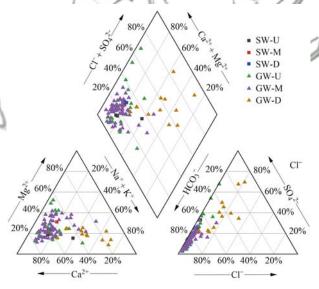
3 结果与分析

3.1 地表水和地下水水化学特征

大通河流域地表水为弱碱性水,地下水为中性偏碱性水(表1). 地表水各河段均以 HCO₃-Mg·Ca型水为主,大通河流域地下水上、中游以 HCO₃-Mg·Ca型水为主,下游以 Cl·SO₄-Na型水为主(图 2). 沿地下水流向,上游地下水中 Mg²⁺浓度随白云石和高镁方解石的溶解而增加,而 Ca²⁺浓度保持相对稳定. 在门源盆地东南侧进入碳酸盐岩的深埋区,由于径流不畅,水中石膏溶解已达到饱和,并出现盐岩溶解(Na⁺和 Cl⁻升高)的现象,SO²⁻浓度随着石膏的溶解而增加. 最后逐渐过渡到碎屑岩地层,由于水位升高,蒸发作用增强^[20],发生水岩作用导致水中

Na⁺浓度升高,而钙镁离子沉淀. 同时,上游地表水 H148 中 ρ (F⁻)为 0. 13 mg·L⁻¹,与其周边潜水 J070 中 F⁻浓度相近;中游地表水 H160 的 F⁻浓度与邻近地下水 J017 和下游 J156、J015 相近. 上述表明,研究区地表水与地下水具有相似的水化学特征,水力联系密切[32].

研究区地表水 TDS 和 EC 值分布较为集中;地下水 TDS 和 EC 值分布较离散,具有明显的径流分区特征;地下水中 TDS 值明显高于地表水[图 3 (a)].地下水在径流过程中,不断地与土壤和围岩中的溶解性盐类发生离子交换,且在蒸发浓缩作用下,地下水中的 TDS 逐渐升高^[33].根据电导特性可推测地下水在河道中的停留时间^[34];根据 TDS[图 3(a)]和 EC 值[图 3(b)]在流域内的分布特征,可以推断出流域内地表水与地下水的转换关系.各河段地下水 TDS 和 EC 不是单调增加的,而在中游降低,中游地下水可能受到 TDS 和 EC 值较低的河水的补给.



SW和GW分别表示地表水和地下水; U、M和D分别表示上、中和下游

图 2 地表水和地下水 Piper 图

Fig. 2 Piper plots of surface water and groundwater

3.2 水化学组分控制因素和来源解析

3.2.1 水化学组分控制因素分析

基于 Gibbs 图分析了各区地表水和地下水针对岩石风化、降水和蒸发浓缩这 3 个控制因子的分布特征(图 4)^[35]. 结果显示, 地表水当量浓度比分布比较集中, 主要落在岩石风化区; 地下水分散分布在岩石风化区和蒸发浓缩区, 其中, 下游地下水大部分落在蒸发浓缩区, 表明受蒸发浓缩因素影响较大, 可能因地下水水位抬升, 地下水排泄地表, 出露地表的泉受到地表的蒸发作用. 部分地下水点落在控制

表 1	地表水和地下水水化学指标特征 ¹⁾	

指标	地表水(N=27)		地下水			
			井水(N=62)		泉水(N=33)	
	范围	均值	范围	均值	范围	均值
EC	302. 0 ~ 599. 0	442. 9	181. 0 ~ 8 180. 0	1 080. 1	342. 0 ~ 4 070. 0	852. 1
pН	8. 0 ~ 8. 7	8.5	5. 6 ~ 8. 4	7. 5	6.0 ~ 8.2	7. 6
TH	143. 8 ~ 283. 7	219.7	193. 7 ~ 1 821. 0	392. 2	197. 0 ~ 3 068. 0	553. 1
TDS	172. 0 ~ 355. 0	279.6	216. 0 ~ 6 180. 0	638. 5	181. 3 ~ 1 665. 0	413.7
K +	1. 24 ~ 2. 54	1.8	0. 9 ~ 26. 8	4. 2	0.9 ~ 8.8	2. 6
Na +	8. 4 ~ 34. 6	13.8	4. 2 ~ 1 394. 0	75.6	4. 4 ~ 390. 1	38. 1
Ca ^{2 +}	39.7 ~61.6	55.8	36. 8 ~ 285. 4	88.6	35. 7 ~ 324. 3	88. 4
Mg^{2+}	9. 8 ~ 34. 5	19. 5	11. 5 ~ 268. 8	41.5	14. 4 ~ 207. 7	46. 8
Cl -	5. 6 ~ 12. 3	8.5	2. 1 ~ 896. 3	60. 9	2. 1 ~ 360. 6	28. 9
SO ₄ ²⁻	21. 3 ~ 89. 3	64. 5	14. 2 ~ 2 980. 0	163.8	6. 1 ~ 1 470. 0	144. 5
HCO_3^-	151. 0 ~ 257. 3	202. 2	143. 8 ~ 688. 6	330.8	180. 6 ~ 961. 1	331. 8
NO_3^-	0.4 ~4.4	0.9	ND. ~26.1	8.4	0. 3 ~ 52. 9	5. 4
F -	0.1 ~ 0.2	0.1	0.1 ~ 3.0	0.4	0.1 ~ 0.4	0. 2

1)单位说明:EC 为 μS·cm $^{-1}$, pH 无量纲,其余指标单位为mg·L $^{-1}$; ND. 表示未检出

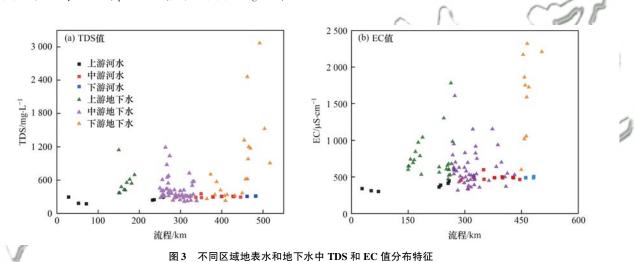


Fig. 3 Distribution characteristics of TDS and EC values of surface water and groundwater in different regions

区虚线外,可能与人类活动、阳离子交换作用等因素有关^[36]. 研究区内地表水和地下水的离子主要由岩体的自然风化和蒸发浓缩引起,离子来源相似,表明二者之间存在某种转换关系.

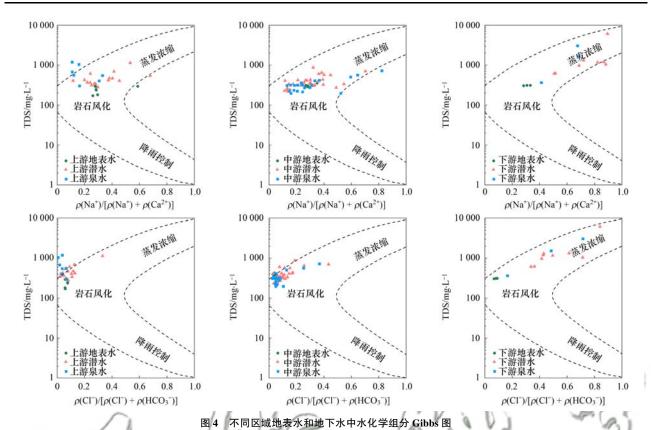
3.2.2 水化学来源解析

通过对研究区 119 组水样 17 个水化学指标的主成分分析,检验结果得到 KMO 值接近 0.70, Bartlett 球形检验量为2 964. 49(P < 0.001),适宜做主成分分析^[37]. 本次共提取 4 个特征值大于 1 的公因子(PC1、PC2、PC3 和 PC4),累计方差贡献率为80. 47%. 其中,水化学指标特征值大于 0.60 的 3 个主成分占总荷载的 76. 50%.

据成分矩阵可知(表 2),主成分 1(PC1)的主要 载荷指标为 EC、TH、TDS、 Na^+ 、 Ca^{2+} 、 Mg^{2+} 、 Cl^- 、 SO_4^{2-} 和 NO_3^- 这 9 项指标,均呈正相关. 研究 区地层岩性以第四系松散沉积岩为主^[15],水岩作用 较强,溶解和迁移的源生矿物,如盐岩(Na^+ 、 Cl^-)、 石膏(Ca^{2+} 、 SO_4^{2-})和白云石(Ca^{2+} 、 Mg^{2+})等,导致水中离子浓度增加,并使 EC、TH 和 TDS 发生变化^[38]. 在氧化-还原反应和阳离子交换作用下,水体中 NH_4^+ 和 NO_2^- 向 NO_3^- 转变. 此外,由于强烈的蒸发作用,使地表水和地下水中的离子富集. 故将 PC1 定义为溶滤-迁移-富集因子.

主成分 2(PC2)的主要载荷指标为高锰酸盐指数和 NH₄*,均呈正相关. 高锰酸盐指数是反映有机物含量特征参数,NH₄* 属于人为活动输入指标^[38],可能由于生活污水和工业废水排入地表水中,在地表水和地下水交换作用下,有机物和含氮污染物渗入地下,使地下水中高锰酸盐指数和 NH₄* 浓度升高. 故将 PC2 定义为人为活动因子.

主成分3(PC3)的主要载荷指标为 As 和 F⁻,均 呈正相关. 含氟矿物在碱性条件下的溶解为 F⁻的富 集提供了条件,硫化物氧化、铁氧化物还原是 As 的 主要来源,而研究区存在含硫、铁等多金属硫化物



Gibbs graphs of water chemistry composition in surface water and groundwater in different regions

表 2 研究区地表水和地下水化学主要指标主成分分析¹⁾
Table 2 Principal component analysis of main chemical indexes

of surface water and groundwater in the study area

or su	mace water and gre	Junuwater n	i the study a	irea		
水化学指标	10 8/7	主成分				
小化子油小	PC1	PC2	PC3	PC4		
EC / p	0. 979	0. 112	0.047	-0.057		
高锰酸盐指数	-0.172	0.636	-0.352	0. 543		
рН	-0.434	0.506	0. 278	-0.397		
TH	0. 939	-0.083	-0.245	-0.066		
TDS	0. 980	0. 115	-0.004	-0.106		
K +	0. 587	0. 324	-0.186	0. 233		
Na +	0. 932	0. 206	0. 141	-0.107		
Ca ^{2 +}	0. 795	-0.246	-0.368	0.049		
Mg^{2} +	0. 943	0.062	-0.113	-0.152		
Cl -	0. 930	0. 182	0.033	-0.143		
SO_4^{2} -	0. 931	0. 199	0.060	-0.174		
HCO ₃	0. 344	-0.559	-0.362	0.424		
NH ₄ ⁺	-0.159	0.724	-0.211	0.432		
NO ₃	0. 615	0.009	-0.138	-0.108		
NO_2^-	-0.122	0.402	0.081	-0.307		
F -	0. 536	-0.014	0.673	0.466		
As	0. 499	-0.115	0.729	0.416		
特征值	8. 559	1. 989	1.657	1.475		
方差贡献率/%	50. 348	11.700	9. 748	8. 676		
累计贡献率/%	50. 348	62.047	71. 795	80. 471		

1)黑体字表示对该因素影响较大的化学指标

矿床^[39]. 煤炭工业基地建在流域中上游, 煤炭中含As 和 F⁻ 矿物经溶滤作用进入地表水和地下水^[40,41]. 故将 PC3 定义为原生沉积环境因子.

综上所述,溶滤-迁移-富集因子(PC1)、人为活动因子(PC2)和原生沉积环境因子(PC3)对研究区地表水和地下水离子的平均贡献率分别为39.1%、15.0%和12.6%,阳离子交替吸附(PC4)和其他离子平均贡献率分别为13.8%和19.5%.结果与Gibbs图得出的主要控制因素一致.

4 讨论

由于本文未采集降水样,在考虑降雨频次、观测距离等条件的情况下,以刘芳等 $^{[42]}$ 所确定的 3 组大气降水点作为参考. 不同水体 δ^{18} O值和 δ D 值变化范围较大(表 3), δ^{18} O值变化幅度为 5. 62‰,标准差为 0. 61; δ D 值变化幅度为 30. 01‰,标准差为 4. 85. 由此看出, δ^{18} O值较 δ D 值更稳定,因此以水体 δ^{18} O值进行补给来源分析. 所有水体d-excess均值为

表 3 大通河流域稳定氘氧同位素取值/‰

Table 3 Values of stable hydrogen and oxygen isotopes in Datong River Basin/%

Tuble 5 values of stable hydrogen and onlygen lestopes in Battery Passas 100							
类型 -	$\delta^{\scriptscriptstyle 1}$	$\delta^{18}{ m O}$		$\delta \mathrm{D}$		d-excess	
	均值 ± 标准差	取值范围	均值 ± 标准差	取值范围	均值 ± 标准差	取值范围	
降水	-6.92 ± 0.41	−7. 34 ~ −6. 52	-42.73 ± 6.45	-49. 99 ~ -37. 63	12. 64 ± 3.36	8. 76 ~ 14. 62	
地下水	-7.86 ± 1.08	$-9.82 \sim -4.2$	-48.4 ± 5.91	−60. 90 ~ −30. 89	14.45 ± 3.43	2. 71 ~ 18. 06	
河水	-7.57 ± 0.33	$-8.07 \sim -7.04$	-46.38 ± 2.18	-50.01 ~ -43.23	14. 16 ± 0.83	12. 73 ~ 15. 90	

13. 75%,大于全球均值(10%),表明夏季流域受强烈的蒸发作用[43].

4.1 同位素端元混合模型

河水 δ^{18} O_{SW}均值(-7.57‰)介于地下水 δ^{18} O_{GW} (-7.86%) 值和降水 $\delta^{18}O_{PW}(-6.92\%)$ 值之间,表 现河水由降水和地下水混合补给特征,故利用二端 元混合模型计算河水不同补给源的贡献率. 结果显 示,流域上游降水贡献率为77.57%,地下水贡献率 为 22. 43%. 此河段地下水同位素较贫化(δ^{18} O $_{GW}$ = -8.72%),降水同位素较富集(δ^{18} O_{PW} = -6.9%), 因此,该河段河水 δ^{18} O_{sw}值较高的主要原因是受到稳 定同位素值较大的降水补给,且降水补给比例也较 大. 中游降水量占 73. 46%, 地下水占 26. 54%, 作为 该河段河水的主要补给源,降水对河水补给速率快, 使河水 δ^{18} O_{sw}值较高. 下游降水和地下水对河水的补 给比例分别为 38.91% 和 61.09%,此外,与中游相 比,下游地下水的稳定同位素值水平降低(δ^{18} O_{GW} = - 8.31‰), 降水同位素值升高(δ^{18} O_{PW} = -6.52%),而河流 δ^{18} O_{sw}值下降,故下游地下水是 河流主要的补给源.

4.2 稳定同位素特征及沿程变化

大通河河水氘氧同位素关系线斜率小于当地大气降水线(LMWL)和全球大气降水线(GMWL),表明河流不仅受大气降水的补给,且受蒸发影响(图5).河水与地下水大多位于降水线方程上方,这是因为降水的补给作用,虽会受到蒸发的影响,但其补给速度更快.流域河水点与地下水点沿 LMWL 两侧分布,表明地表水和地下水总补给来源为当地降水.除受降水补给外,流域河水氘氧同位素关系线更靠

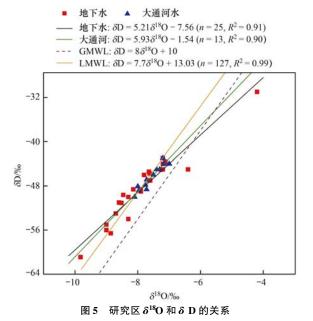


Fig. 5 Relationship between δ^{18} O and δ D in the study area

近地下水线,显示河水更易受地下水影响.

河水自上而下流动的过程中,由于受其他水源 补给和环境等因素的影响,造成了不同的补排关系, 从而使河流中的稳定氘氧同位素值发生了变化. 由 图 6(a) 可知,不同河段河水 δ^{18} O_{sw} 值随着流程(至 河源的距离)的增大而有显著差异:流域 δ^{18} O_{sw}值以 点 H028 为界,先增大,随后以点 H151 为界起伏减 小,其 δ^{18} O_{sw}值呈富集到贫化的变化趋势. 上游点 H32~H28 受偏正降水δ¹δОρw值的补给较大,而靠近 下游的点 H149 受贫化地下水(δ^{18} O_{GW}均值 – 7.7‰) 补给较大,结果表明河水最初补给源为降水,流动过 程中经历了蒸发作用,最后因受地下水补给, δ^{18} O_{sw} 呈贫化特征. 此外,H151 为纳子峡水库水,与库前河 水相比,水库水由于蒸发作用δ¹⁸O_{sw}同位素富集,由 TDS 计算得出蒸发损失可达 2.7%~14.8%, 故水库 蓄水会导致地表水分的大量蒸发,在干旱地区是一 种极大的水资源浪费. 下游低海拔区地势平坦开阔, 地下水以泉的方式补给河水,但由于水电站蓄水致 使区内部分河段断流[4],河水与地下水之间转化较 弱[图6(c)].

由图 6(b)可知,大通河流域河水 d-excess 值随海拔升高而波动降低的特点,采样时间段为夏季(6~8月),该区河水 d-excess 值均高于全球雨水 d-

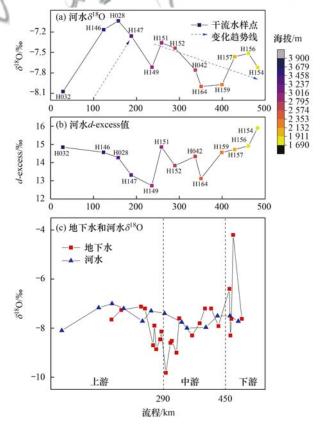


图 6 稳定同位素沿流程变化

Fig. 6 Variation in stable isotopes along the distance

excess 均值(10‰),夏季气温较高,河流蒸发强烈, 因此,当地蒸发的水汽会对降水补给产生一定的影响.此外,由于河流自高海拔向低海拔区流动,地下水与河流间存在着紧密的补排关系,从而导致了 d-excess 值在沿程发生了明显的变化.

5 地表水与地下水的转化关系

大通河上游分布海拔3 000 m以上的高山,区内 冰雪融水、降水和基岩裂隙水形成的地表径流在山 前补给地下水, 氘氧稳定同位素结果表明降水是河 段地表水和地下水总的补给来源,补给比例为 77.57%. 而中游地下水 TDS 和 EC 值沿程减小,表 明地下水受到了地表水的补给稀释,这与二端元模 型结果一致,河段地表水和地下水均为 HCO₃-Mg· Ca型. 冲洪积扇区降水和基岩裂隙水入渗补给地下 水,由于中新世中晚期祁连山地区进入了隆升期,盆 山高差较大[45],山前地下水因构造影响,径流受阻, 地下水位升高,以泉的形式溢出补给河水,使河水 TDS 值增加, 氘氧同位素值贫化. 下游由于水电站蓄 水使区内部分河段断流,河水与地下水之间转化较 弱,河水为 HCO,-Mg·Ca 型,地下水为 Cl·SO,-Na 型.此外,随着路径的增长,河水和地下水的蒸发浓 缩作用强烈,河水在下游蒸发比例达 2.7%~ 14.8%.根据二端元模型结果显示,流域上、中游地 区河水主要受大气降水补给,平均补给比例分别为 77.57%和73.46%;下游河段主要为潜水和泉水溢 出补给河水,平均补给比例为61.09%.河流自上游 至下游,降雨补给比例减小,地下水补给比例增加, 除受构造运动影响外,可能与海拔变化有关,高海拔 山区地下水埋深大于海拔低的地区,地下水在高海 拔区更易形成山间径流,低海拔区易溢出地表.根据 地表水和地下水的水文地球化学特征可知,上、中 游为地表水补给径流区,下游为地下水的排泄区.

6 结论

- (1) 大通河流域地表水为偏碱性淡水,以HCO₃-Mg·Ca型为主; 地下水为偏碱性微咸水,上游和中游主要为 HCO₃-Mg·Ca型,下游为 Cl·SO₄-Na型.溶滤-迁移-富集因子、人类活动因子和原生地质因子对地表水和地下水中离子的平均贡献率分别为 39.1%、15.0%、12.6%.上游地下水存在少量钙、镁长石的溶解,中游地表水和地下水主要来自碳酸盐岩的溶解,下游地下水消耗有机物产生较多的 CO₂ 溶解到水里使碳酸盐含量升高,导致碳酸钙较多的溶解.
 - (2)大通河河水氘氧同位素含量沿地下水流向

呈富集到贫化的变化趋势,中上游 δD 和δ¹⁸O较下游 富集;大通河上、中游地区河水主要受大气降水补 给,平均补给比例分别为 77.57% 和 73.46%,均为 地表水补给区;下游河段受地质构造和水文地质条 件等影响,主要为潜水和泉水溢出补给河水,平均补 给比例为 61.09%,为地下水排泄区.

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